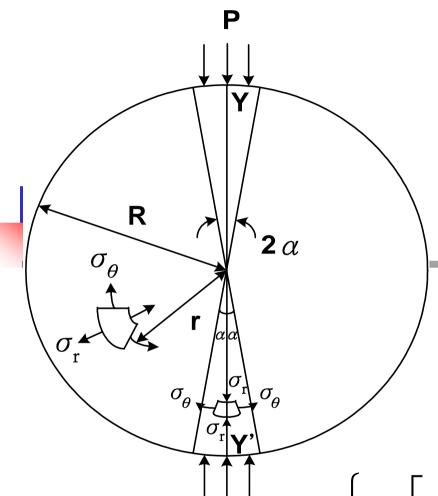
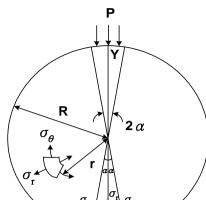
#### **Brazilian Test**

(Diametral Compression of Discs)



The stress component normal to the loading diameter YY' ( $\sigma_{\theta}$ ) (Hondros, 1959)

$$\sigma_{\theta} = -\frac{P}{\pi R t \alpha} \left\{ \frac{\left[1 - (\frac{r}{R})^{2}\right] \sin 2\alpha}{1 - 2(\frac{r}{R})^{2} \cos 2\alpha + (\frac{r}{R})^{4}} - \tan^{-1} \left[\frac{1 + (\frac{r}{R})^{2}}{1 - (\frac{r}{R})^{2}} \tan \alpha\right] \right\}$$



The stress component along YY'  $(\sigma_r)$ 

$$\frac{P}{\pi R t \alpha} \left\{ \frac{\left[1 - \left(\frac{r}{R}\right)^{2}\right] \sin 2\alpha}{1 - 2\left(\frac{r}{R}\right)^{2} \cos 2\alpha + \left(\frac{r}{R}\right)^{4}} + \tan^{-1} \left[\frac{1 + \left(\frac{r}{R}\right)^{2}}{1 - \left(\frac{r}{R}\right)^{2}} \tan \alpha\right] \right\}$$

where  $\sigma_{\theta}$ : stress component normal to the loading diameter

 $\sigma_r$ : stress component along the loading diameter

P: applied force

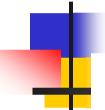
r: distance from the center of disc

t: thickness of the disc

 $2\alpha$ : angular distance over which P is assumed to be distributed radially

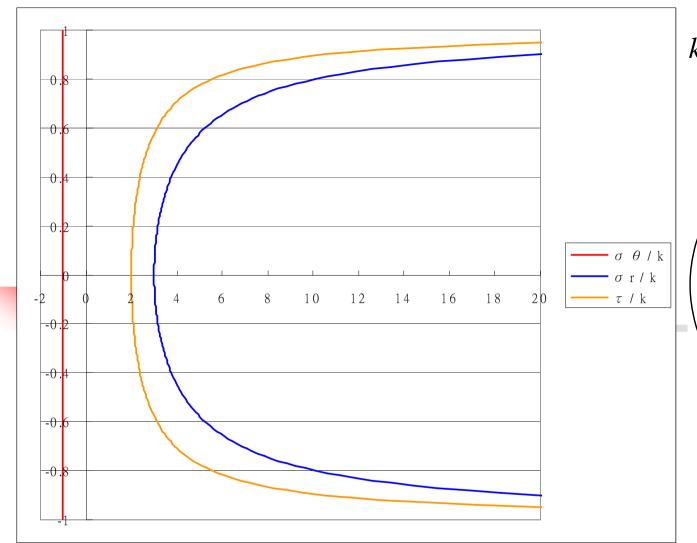
R: radius of the disc

 $\sigma_{\theta}$  及  $\sigma_{r}$  差別在  $\tan^{-1}$  前的加減號及最前面的正、負號



$$\sigma_{\theta} = -\frac{P}{\pi R t \alpha} \left\{ \frac{\left[1 - \left(\frac{r}{R}\right)^{2}\right] \sin 2\alpha}{1 - 2\left(\frac{r}{R}\right)^{2} \cos 2\alpha + \left(\frac{r}{R}\right)^{4}} - \tan^{-1} \left[\frac{1 + \left(\frac{r}{R}\right)^{2}}{1 - \left(\frac{r}{R}\right)^{2}} \tan \alpha\right] \right\}$$

$$\sigma_{r} = \frac{P}{\pi R t \alpha} \left\{ \frac{\left[1 - (\frac{r}{R})^{2}\right] \sin 2\alpha}{1 - 2(\frac{r}{R})^{2} \cos 2\alpha + (\frac{r}{R})^{4}} + \tan^{-1} \left[\frac{1 + (\frac{r}{R})^{2}}{1 - (\frac{r}{R})^{2}} \tan \alpha\right] \right\}$$



$$\sigma_{\theta} = -\frac{P}{\pi R t \alpha} \left\{ \frac{\left[1 - (\frac{r}{R})^{2}\right] \sin 2\alpha}{1 - 2(\frac{r}{R})^{2} \cos 2\alpha + (\frac{r}{R})^{4}} - \tan^{-1} \left[\frac{1 + (\frac{r}{R})^{2}}{1 - (\frac{r}{R})^{2}} \tan \alpha\right] \right\} \quad \sigma_{r} = \frac{P}{\pi R t \alpha} \left\{ \frac{\left[1 - (\frac{r}{R})^{2}\right] \sin 2\alpha}{1 - 2(\frac{r}{R})^{2} \cos 2\alpha + (\frac{r}{R})^{4}} + \tan^{-1} \left[\frac{1 + (\frac{r}{R})^{2}}{1 - (\frac{r}{R})^{2}} \tan \alpha\right] \right\} \quad 102$$

當試體承受線荷重(即  $\alpha$  很小),則在圓心處(r=0)張力最大

If 
$$r = 0$$

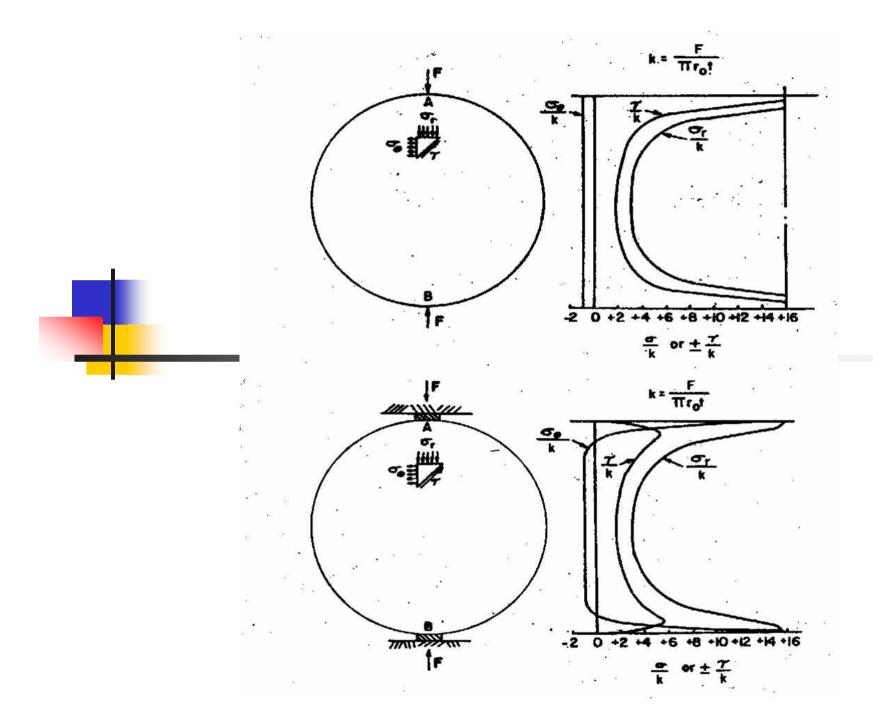
$$\sigma_{\theta} = -\frac{P}{\pi R t} \left[ \frac{\sin 2\alpha}{\alpha} - 1 \right] \approx -\frac{P}{\pi R t} = \sigma_{t} = -\frac{2P}{\pi D t} = \frac{0.636 P}{Dt}$$

$$\sigma_r = \frac{P}{\pi R t} \left[ \frac{\sin 2\alpha}{\alpha} + 1 \right] \approx \frac{3P}{\pi R t} = \frac{6P}{\pi D t} = -3\sigma_t$$

在圓心處壓力爲張力的 $\mathbf{3}$ 倍,但  $\sigma_c/\sigma_t$ 一般在 $\mathbf{8}\sim\mathbf{50}$ ,故將產生張力破壞



Bazilian test has been found to give tensile strength higher than that of direct tension test, probably owing to the effect of fissures. Short fissures weaken a direct tension specimen more severely than they weaken a splitting tension specimen. The ratio has been found to vary from 1 to more than ten as length of preexisting fissures grows larger.



1.試簡述下列各試驗之方法與目的:(88檢覈)

(一)岩石巴西人法試驗(Brazilian test)

(註:必要時,請輔以簡圖說明之)

- 2.請列舉數種岩石試體張力強度之試驗方法,並比較各種方法之優劣。(15分)(8專技)(85應地)
- 3.在實驗室內,求取岩石材料張力強度的方法有那些?請說 明其求法?(15分)

1.若有一圓板,直徑爲 d,厚I,在圓板上下端施加集中荷重 P

,則圓板內部任意點A處所受應力爲:

$$\sigma_{x} = \frac{2P}{\pi \ell} \left( \frac{\sin^{2}\theta_{1}\cos\theta_{1}}{r_{1}} + \frac{\sin^{2}\theta_{2}\cos\theta_{2}}{r_{2}} \right) - \frac{2P}{\pi d\ell}$$

$$\sigma_{y} = \frac{2P}{\pi \ell} \left( \frac{\cos^{3}\theta_{1}}{r_{1}} + \frac{\cos^{3}\theta_{2}}{r_{2}} \right) - \frac{2P}{\pi d\ell}$$

$$\tau_{xy} = \frac{2P}{\pi \ell} \left( \frac{\cos^{2}\theta_{1}\sin\theta_{1}}{r_{1}} + \frac{\cos^{2}\theta_{2}\sin\theta_{2}}{r_{2}} \right)$$

依此原理,製作圓板形岩石試體,進行壓製試驗,又稱巴西試驗(BrazilianTest),請繪出a,b間之 $\sigma_x$ , $\sigma_y$ , $\tau_{xy}$ 的變化,並說明應如何求試體的張力強度。**(20**分**)** 

1.解釋或說明下列各項:(91專技)

說明完整岩心在進行巴西抗張試驗時,圓盤試體中心位置之 應力狀態。(5分)

- 1.有數公尺甚爲完整,不具不連續面,且具有相同岩性之岩心
  - ,其直徑為**50mm**,平均岩石密度為**0.025MN/m³**,經以下一系列之試驗:
  - (一)徑向點荷重試驗,得平均破壞荷重為5KN(Kilo-Newton)
  - (二)單壓試驗,得平均破壞荷重為100KN
  - (三)巴西試驗,得平均破壞荷重為10KN(厚度為直徑一半)
  - (四)超音波試驗,得平均縱波波速為3500m/sec,平均橫波 波速為2000m/sec

#### 請回答:

- (一)本岩石之張力強度爲何?(5分)
- (二)本岩石動態彈性係數爲何?(10分)
- (三)未來類似之岩石如何由徑向點荷重試驗結果估計其單壓強度?(5分)

1.以直徑5公分之乾燥岩心進行試驗,得結果如下:

(一)巴西劈裂試驗:岩心直徑5公分,厚度3公分,劈裂荷重 = 236 kgf

(二)一軸壓縮試驗:岩心直徑5公分,高度10公分,壓縮破壞荷重 =

1570kgf

(三)三軸壓縮試驗:岩心直徑5公分,高度10公分,試驗結果:

圖 壓(kgf/cm <sup>2</sup> )	破壞時之軸差荷重	(kgf)
20	2160 2900	
80	3965	

請將此三種試驗的破壞莫爾(Mohr)應力圓繪出,並繪製莫爾一庫倫破壞 準則(Mohr-Coulomb criterion),以及預測圍壓爲70 kgf/cm²時,岩心試 體破壞強度和破壞面之角度。(25分)

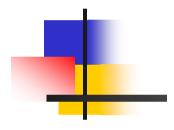
#### Read

- (1).Vutukuri, Lama and Saluja, vol.1, p.15~111 (1974).
- (2).Jaeger and Cook, "Fundamental of Rock Mechanics" p.169~170 and p.258~260 (1979).
  - (3).Brown, "ISRM Suggested Methods", p.120~121 (1981).
  - (4).ASTM D3967-86 "Standard Method for Splitting Tensile Strength of Intact Rock Core Specimens".

4.4 Behaviour of isotropic rock material in multiaxial compression

4.4. I Types of multiaxial compression test

Biaxial compression (  $\sigma_1 \ge \sigma_2$  ,  $\sigma_3 = 0$  )



Triaxial compression (  $\sigma_1 > \sigma_2 = \sigma_3$  ) Conventional Triaxial

Polyaxial compression ( $\sigma_1 > \sigma_2 > \sigma_3$ ) True Triaxial

# 4.4.2 Biaxial compression ( $\sigma_1 \ge \sigma_2$ , $\sigma_3 = 0$ )

Biaxial compression tests are carried out by applying different normal stresses to two pairs of faces of a cube, plate or rectangular prism of rock.

# 4.4.2 Biaxial compression ( $\sigma_1 \ge \sigma_2$ , $\sigma_3 = 0$ )

The great difficulty with such tests is that the end effects described in section 4.3.3 exert an even greater influence on the stress distribution within the specimen than in the case of uniaxial compression. For

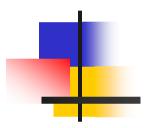
this reason, fluid rather than solid medium loading is preferred.

### 4.4.2 Biaxial compression ( $\sigma_1 \ge \sigma_2$ , $\sigma_3 = 0$ )

Brown (1974) carried out a series of biaxial compression tests on 76 mm square by 25 mm thick plates of Wombeyan Marble which were loaded on their smaller faces through (a) 76 mm x 25 mm solid steel platens, and (b) brush platens made from 3.2 mm square steel pins.

### 4.4.2 Biaxial compression ( $\sigma_1 \ge \sigma_2$ , $\sigma_3 = 0$ )

Figure 4.15 shows the peak strength envelopes obtained in tests carried out at constant ratios. The data are normalised with respect to the uniaxial compressive strength of the plates  $\sigma_c = 66$  MPa.



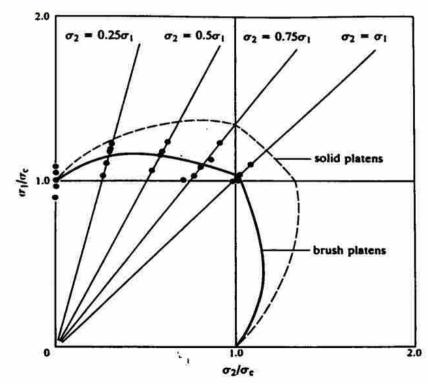


Figure 4.15 Biaxial compression test results for Wombeyan Marble (after Brown, 1974).

### 4.4.2 Biaxial compression ( $\sigma_1 \ge \sigma_2$ , $\sigma_3 = 0$ )

This was attributed to the influence of end effects. When the

brush platens were used, the maximum measured increase in peak strength over  $\sigma_c$  was only 15%. For  $\sigma_2 = \sigma_1$  , no strength increase. The practical consequence of these results is that, for this rock type, the 'strengthening' effect of the intermediate principal stress can be neglected so that the uniaxial compressive strength,  $\sigma_c$  , should be used as the rock material strength whenever  $\sigma_3 = 0$ . This slightly conservative conclusion is likely to apply to a wide range of rock types.

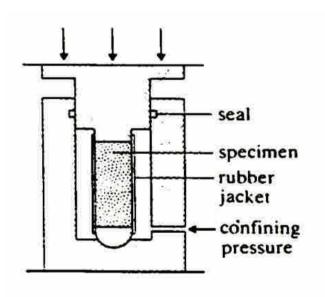
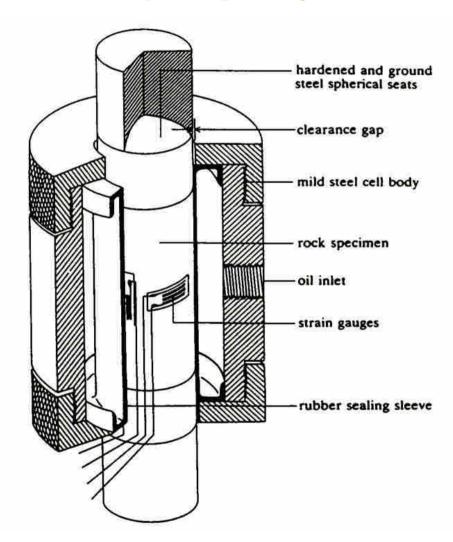


Figure 4.16 Elements of a conventional triaxial testing apparatus.

The specimen is placed inside a pressure vessel (Figures 4.16 and 4.17) and a fluid pressure,  $\sigma_3$ , is applied to its surface. A jacket, usually made of a rubber compound, is used to isolate the specimen from the confining fluid which is usually oil. The axial stress,  $\sigma_1$ , is applied to the specimen via a ram passing through a bush in the top of the cell and hardened steel end caps.



Figure 4.17 Cut-away view of the triaxial cell designed by Hoek and Franklin (1968). Because this cell does not require drainage between tests, it is well suited to carrying out large numbers of tests quickly.



Axial deformation of the specimen may be most conveniently monitored by linear variable differential transformers (LVDTs) mounted inside or outside the cell, but preferably inside. Local axial and circumferential strains may be measured by electric resistance strain gauges attached to the surface of the specimen (Figure 4.17).

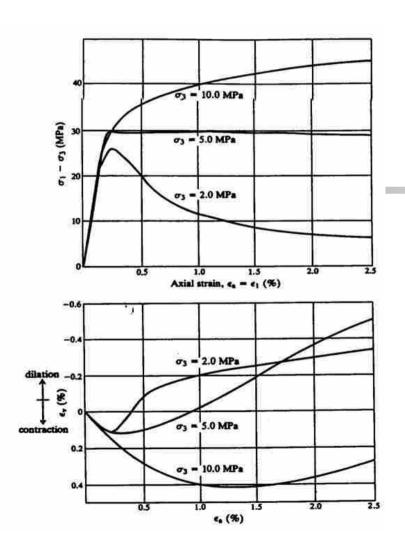
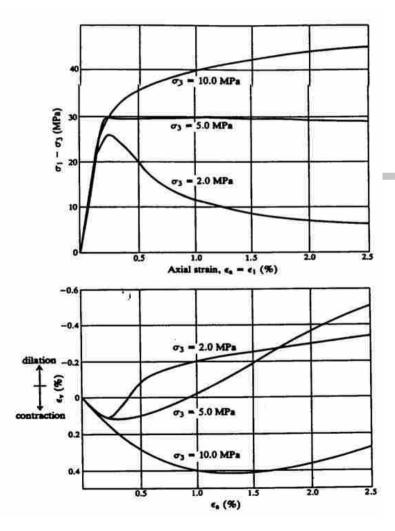


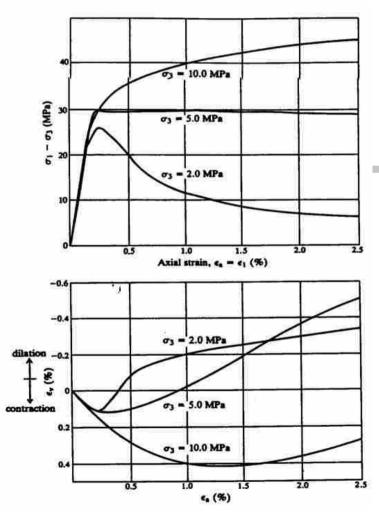
Figure 4.18 shows some results obtained using such a system in tests carried out at three different confining pressures on specimens of an oolitic limestone. An important feature of the behaviour of rock material in triaxial compression is illustrated by Fig4.18.

Figure 4.18 Results of triaxial compression tests on an oolitic limestone with volumetric strain measurement (after Elliott, 1982).



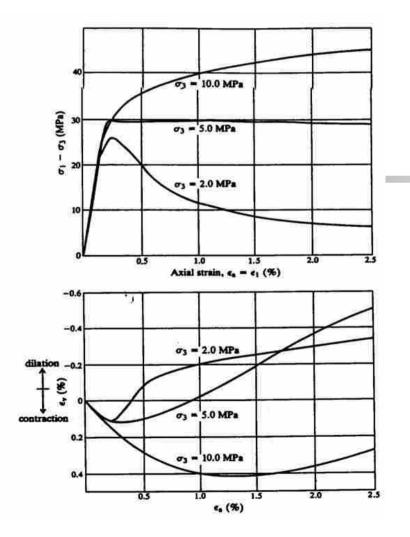
When the specimen is initially loaded it compresses, but a point is soon reached, generally before the peak of the axial stress-axial strain curve, at which the specimen begins to dilate (increase in volume) as a result of internal fracturing.

Figure 4.18 Results of triaxial compression tests on an oolitic limestone with volumetric strain measurement (after Elliott, 1982).



Shortly after the peak strength is reached, the net volumetric strain of the specimen becomes dilational. Dilation continues in the post-peak range. The amount of dilation decreases with increasing confining pressure.

Figure 4.18 Results of triaxial compression tests on an oolitic limestone with volumetric strain measurement (after Elliott, 1982).

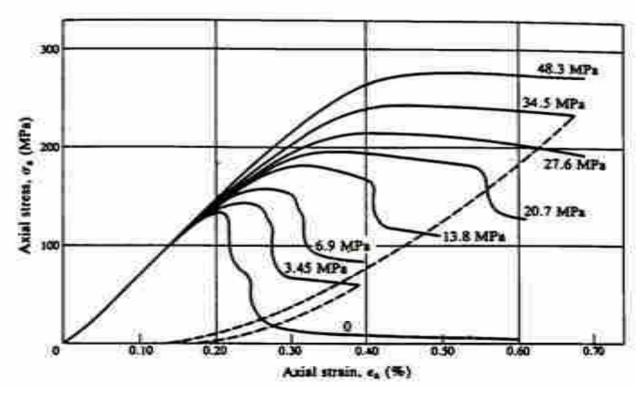


At very high confining pressures, often outside the range of engineering interest, dilation may be totally suppressed with the volumetric strains remaining contractile throughout the test.

Figure 4.18 Results of triaxial compression tests on an oolitic limestone with volumetric strain measurement (after Elliott, 1982).

Figure 4.19 illustrates a number of other important features of the behaviour of rock in triaxial compression.

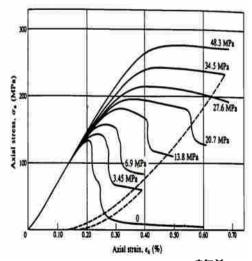
Figure 4.19 Complete axial stress -axial strain curves obtained in triaxial compression tests on Teonessee Marble as the confining pressures indicated by the numbers on the curves (after Wawersik and Fairfurst, 1970).



These and similar data for other rocks show that, with increasing confining pressure,

- (a) the peak strength increases;
- (b) there is a transition from typically brittle to fully ductile behaviour with the introduction of plastic mechanisms of deformation including cataclastic flow and grain-sliding effects;

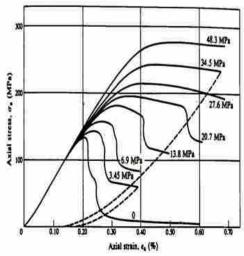
Figure 4.19 Complete axial stress—axial strain curves obtained in trusxial compression tests on Teonessee Marble at the confining pressures indicated by the numbers on the curves (after Wawersik and Fairfurns, 1970).



These and similar data for other rocks show that, with increasing confining pressure,

- (c) the region incorporating the peak of the ( $\sigma_a$   $\varepsilon_a$ ) curve flattens and widens;
- (d) the post-peak drop in stress to the residual strength reduces and disappears at high values of  $\sigma_3$  .





The confining pressure at which the post-peak reduction in strength disappears and the behaviour becomes fully ductile ( $\sigma_3 = 48.3$  MPa in Figure 4.19), is known as the brittle-ductile transition pressure and varies with rock type.

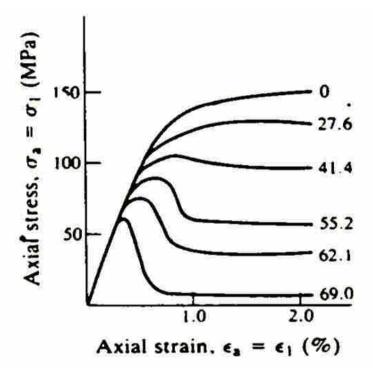


Figure 4.20 Effect of pore pressure (given in MPa by the numbers on the curves) on the stress-strain behaviour of a limestone tested at a constant confining pressure of 69 MPa (after Robinson, 1959). The influence of pore-water pressure

Brace and Martin (1968) conducted triaxial compression tests on a variety of crystalline silicate rocks of low porosity (0.001- 0.03) at axial strain rates of  $10^{-3} - 10^{-8} \, s^{-1}$ . They found that the classical effective stress law held only when the strain rate was less than some critical value which depended on the permeability of the rock, the viscosity of the pore fluid and the specimen geometry. At strain rates higher than the critical, static equilibrium could not be achieved throughout the specimen.

### 4.4.4 Polyaxial compression ( $\sigma_1 > \sigma_2 > \sigma_3$ )

These tests may be carried out on cubes or rectangular prisms of rock with different normal stresses being applied to each pair of opposite faces. The results of polyaxiai compression tests on prismatic specimens are often conflicting, but generally indicate some influence of the intermediate principal stress,  $\sigma_2$ , on stress-strain behaviour.

### 4.4.4 Polyaxial compression ( $\sigma_1 > \sigma_2 > \sigma_3$ )

Generally, the peak strength increases with increasing  $\sigma_2$  for constant  $\sigma_3$ , but the effect is not as great as that caused by increasing  $\sigma_3$  by a similar amount (Paterson, 1978). However, doubts must remain about the uniformity of the applied stresses in these tests and the results should be interpreted with great care.

#### 4.4.5 Influence of stress path

A test of considerable relevance in this regard is the triaxial extension test which is carried out in a triaxial cell with the confining pressure,  $\sigma_r$ , greater than the axial stress,  $\sigma_a$ . The test may be commenced at  $\sigma_a = \sigma_r$  with  $\sigma_a$ , being progressively reduced so that  $\sigma_r = \sigma_1 = \sigma_2 > \sigma_a = \sigma_3$ . With modern servo-controlled testing machines, almost any desired total or effective stress path can be followed within the limitations imposed by the axisymmetric configuration of the triaxial cell.

#### 4.4.5 Influence of stress path

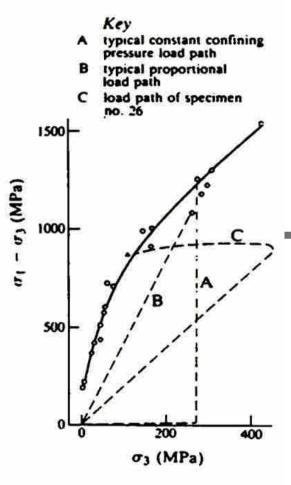


Figure 4.21 Influence of stress path on the peak strength envelope for Westerly Granite (after Swanson and Brown, 1971).

Swanson and Brown (1971) investigated the effect of stress path on the peak strength of a granite and a quartz diorite. They found that, for both rock types, the peak strengths in all tests fell on the same envelope (Figure 4.21 for Westerly Granite) irrespective of stress path. They also found that the onset of dilatancy, described in section 4.4.3, is stress-path independent. Similarly, Elliott (1982) found the yield locus of a high-porosity, oolitic limestone to be stress-path independent.

# 4.5 Strength criteria for isotropic rock material4.5.1 Types of strength criterion

Peak Strength Criterion : A relation between stress components which will permit the peak strengths developed under various stress combinations to be predicted.

Residual Strength Criterion : be used to predict residual strengths under varying stress conditions.

Yield Criterion: A relation between stress components which is satisfied at the onset of permanent deformation.

Strength and yield criteria are best written in effective stress form

# 4.5 Strength criteria for isotropic rock material4.5.1 Types of strength criterion

The general form of the peak strength criterion should be



$$\sigma_1 = f(\sigma_2, \sigma_3)$$

This is sometimes written in terms of the shear and normal stresses, a on a particular plane in the specimen

$$\tau = f(\sigma_n)$$

All of the criteria used in practice are reduced to the form

$$\sigma_1 = f(\sigma_3)$$

$$s = c + \sigma_n \tan \phi \tag{4.11}$$

where c = cohesion and  $\phi = angle$  of internal friction.

$$\sigma_n = \frac{1}{2}(\sigma_1 + \sigma_3) + \frac{1}{2}(\sigma_1 - \sigma_3)\cos 2\beta$$

and

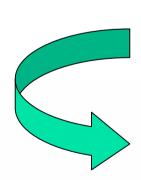
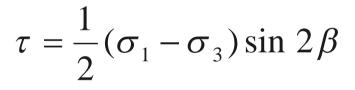
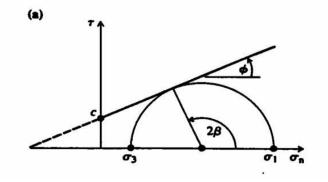
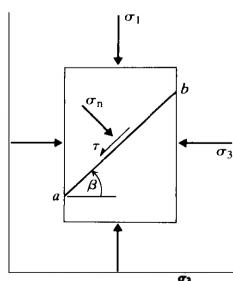


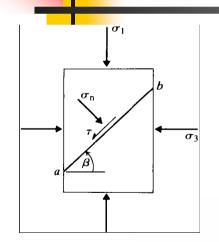
Figure 4.23 Coulomb strength envelopes in terms of (a) shear and normal stresses, and (b) principal stresses.





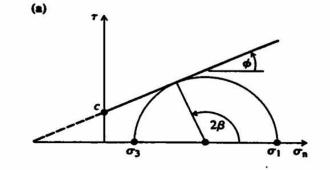


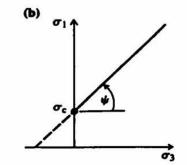
Substitution for  $\sigma_n$  and  $s=\tau$  in equation 4.11 and rearranging gives the limiting stress condition on any plane defined by  $\beta$  as



$$\sigma_1 = \frac{2c + \sigma_3 \left[\sin 2\beta + \tan \phi \left(1 - \cos 2\beta\right)\right]}{\sin 2\beta - \tan \phi \left(1 + \cos 2\beta\right)} \tag{4.12}$$

Figure 4.23 Coulomb strength envelopes in terms of (a) shear and normal stresses, and (b) principal stresses.





$$\sigma_1 = \frac{2c + \sigma_3 \left[\sin 2\beta + \tan \phi \left(1 - \cos 2\beta\right)\right]}{\sin 2\beta - \tan \phi \left(1 + \cos 2\beta\right)}$$

The Mohr circle construction of Figure 4.23a gives the orientation of this critical plane as

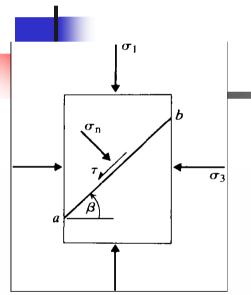
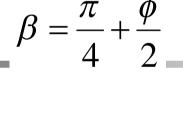
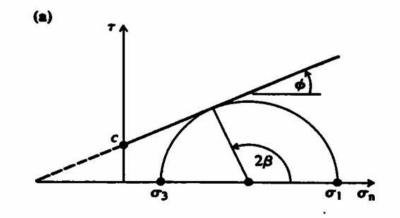
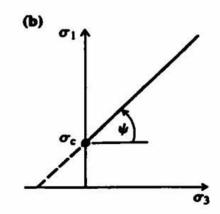


Figure 4.23 Coulomb strength envelopes in terms of (a) shear and normal stresses, and (b) principal stresses.







$$\sigma_1 = \frac{2c + \sigma_3 \left[\sin 2\beta + \tan \phi \left(1 - \cos 2\beta\right)\right]}{\sin 2\beta - \tan \phi \left(1 + \cos 2\beta\right)}$$

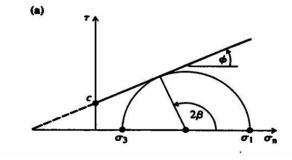
$$\beta = \frac{\pi}{4} + \frac{\phi}{2}$$

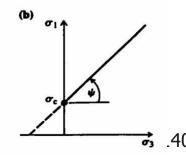
For the critical plane,  $\sin 2\beta = \cos \phi$  ,  $\cos 2\beta = -\sin \phi$  , and equation 4.12 reduces to

$$\sigma_1 = \frac{2c\cos\phi + \sigma_3(1+\sin\phi)}{1-\sin\phi}$$
 (4.14)

$$\sigma_1 = \sigma_3 K_p + 2c \sqrt{K_p}$$

Figure 4.23 Coulomb strength envelopes in terms of (a) shear and normal stresses, and (b) principal stresses.





The **Slope** of (b) linera equation of  $\sigma_1$   $\sigma_3$ 

$$\tan \psi = \frac{1 + \sin \phi}{1 - \sin \phi} = K_{p}$$

# interception

$$\sigma_c = \frac{2c\cos\phi}{1-\sin\phi} = 2c\sqrt{K_p}$$

(4.15)

(4.16)

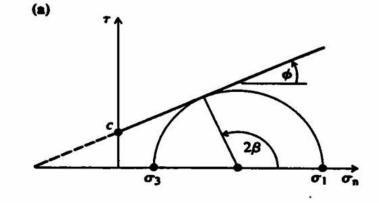
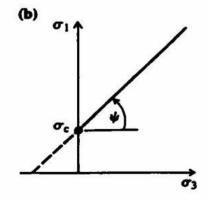


Figure 4.23 Coulomb strength envelopes in terms of (a) shear and normal stresses, and (b) principal stresses.



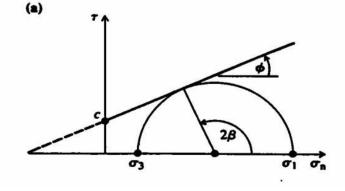
If the Coulomb envelope shown in Figure 4.23b is extrapolated to  $\sigma_1=0$ , it will intersect the  $\sigma_3$  axis at an apparent value of uniaxial tensile strength of the material given by

$$\sigma_T = \frac{2c\cos\phi}{1+\sin\phi}$$

(b)  $\sigma_1$ 

(4.17)

Figure 4.23 Coulomb strength envelopes in terms of (a) shear and normal stresses, and (b) principal stresses.





A tensile cutoff is usually applied at a selected value of uniaxiai tensile stress, as shown in Figure 4.24. For practical purposes, it is prudent to put .

$$T_0 = 0$$

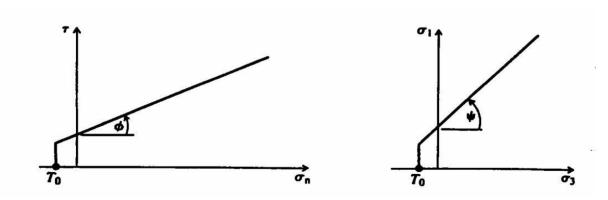
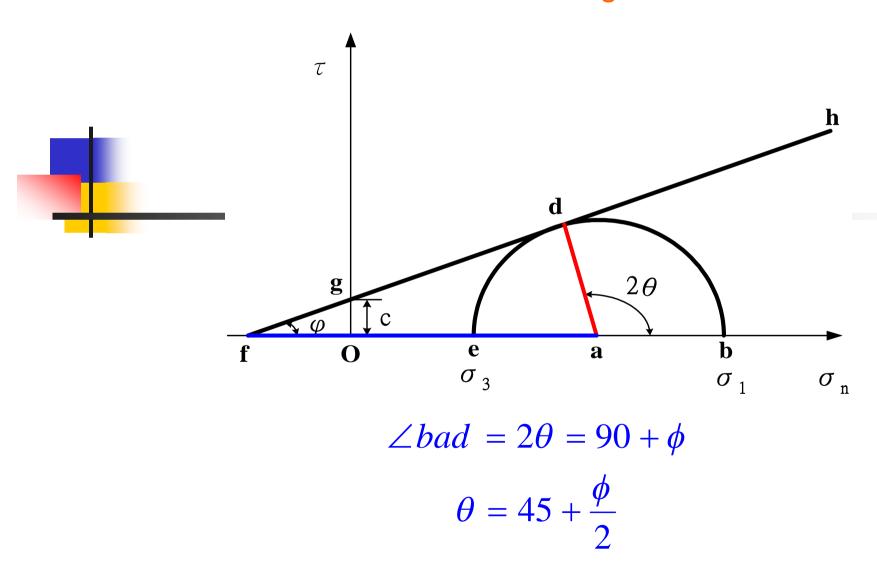
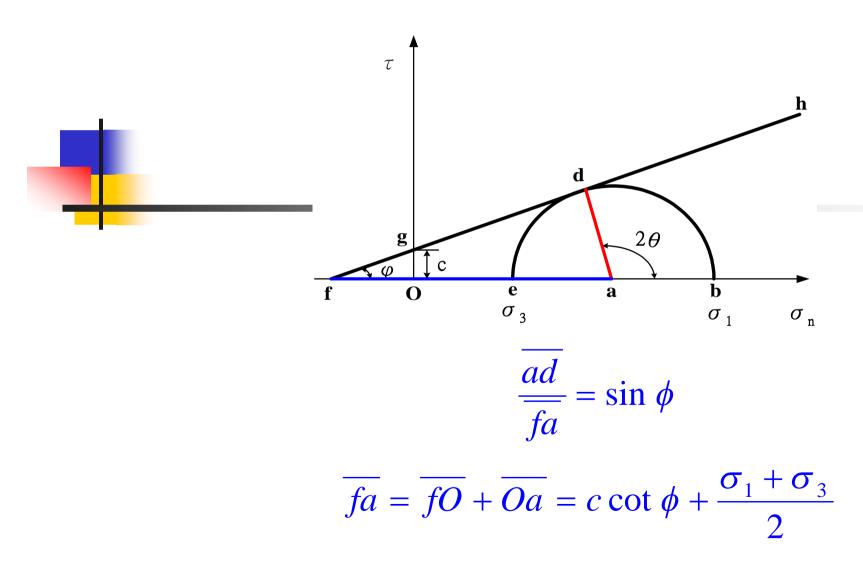


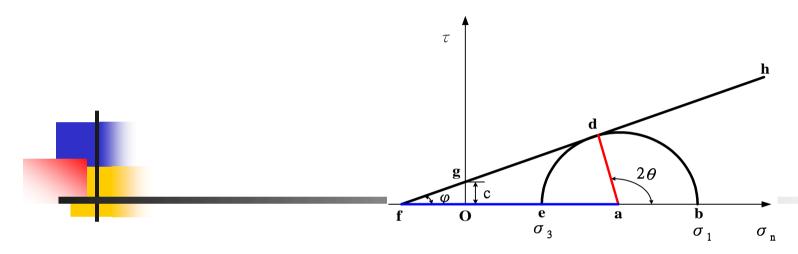
Figure 4.24 Coulomb strength envelopes with a tensile cut-off.

Although it is widely used, Coulomb's criterion is not a particularly satisfactory peak strength criterion for rock material. The reasons for this are:

- (a) It implies that a major shear fracture exists at peak strength. Observations such as those made by Wawersik and Fairhurst (1970) show that this is not always the case.
- (b) It implies a direction of shear failure which does not always agree with experimental observations.
- (c) Experimental peak strength envelopes are generally non-linear. They can be considered linear only over limited ranges of  $\sigma_n$  or  $\sigma_3$ .



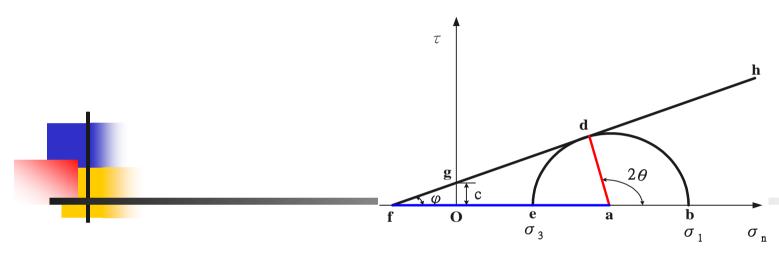




$$\overline{ad} = \frac{\sigma_1 - \sigma_3}{2} \tag{5}$$

Substituting for Eq.(4) and Eq.(5) in Eq.(3) gives

$$\sin \phi = \frac{\frac{\sigma_1 - \sigma_3}{2}}{c \cot \phi + \frac{\sigma_1 + \sigma_3}{2}}$$
 (6)



$$\sigma_1 = \sigma_3 \left( \frac{1 + \sin \phi}{1 - \sin \phi} \right) + 2c \sqrt{\left( \frac{1 + \sin \phi}{1 - \sin \phi} \right)}$$
 (7)

or

$$\sigma_1 = \sigma_3 \tan^2(45 + \frac{\phi}{2}) + 2c \tan(45 + \frac{\phi}{2})$$
 (8)

or

$$\sigma_1 = \sigma_3 K_p + 2c \sqrt{K_p}$$
 (9)

某一岩層,強度受莫耳-庫倫((Mohr-Coulomb)破壞準則控制,其凝聚力C=20MPa,內摩擦角 $\phi=40^{\circ}$ ,今受到最大主應力為100Mpa,最小主應力為15MPa的應力作用。

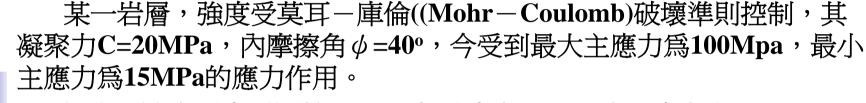
(一)試證明該一岩層在此應力狀況下是穩定的。

故穩定

$$\sigma_{1} = \sigma_{3} \left( \frac{1 + \sin \phi}{1 - \sin \phi} \right) + 2c \sqrt{\left( \frac{1 + \sin \phi}{1 - \sin \phi} \right)} = \sigma_{3} K_{p} + 2c \sqrt{K_{p}}$$

$$= 15 \times 4.6 + 2 \times 20 \times \sqrt{4.6}$$

$$= 154.8 MPa > 100 MPa$$



(二)假設孔隙水壓由零開始上升,直到破壞發生,試問破壞時 之孔隙水壓力大小爲何?

$$(\sigma_1 - U_f) = (\sigma_3 - U_f)K_p + 2c\sqrt{K_p}$$

$$(100 - U_f) = (15 - U_f)4.6 + 2 \times 20 \times \sqrt{4.6}$$

$$3.6 \times U_f = 54.78$$

$$U_f = 15.22 MPa$$

某一岩層,強度受莫耳-庫倫((Mohr-Coulomb)破壞準則控制,其凝聚力C=20MPa,內摩擦角 $\phi=40^{\circ}$ ,今受到最大主應力為100Mpa,最小主應力為15MPa的應力作用。

(三)破壞時,斷裂面上的有效正應力與剪應力分別是多少?

$$\alpha = 45 + \frac{\phi}{2} = 65$$

$$\sigma_n = \frac{\sigma_1 + \sigma_3}{2} - \frac{\sigma_1 - \sigma_3}{2} \cos(\pi - 2 \times 65) = 15MPa$$

$$\tau = \frac{\sigma_1 - \sigma_3}{2}\sin(\pi - 2 \times 65) = 32.56MPa$$